Correlation of dissolved gaseous mercury and radon profiles in the Mediterranean Basin seawater

Janja Vaupotič¹, Jože Kotnik¹, Milena Horvat¹ & Nicola Pirrone²

¹Jožef Stefan Institute, Jamova 39, 1000 Ljubljana, Slovenia; E-mail: janja.vaupotic@ijs.si. ²CNR, Institute for Atmospheric Pollution, Rende, Italy.

Abstract: Vertical profiles of dissolved gaseous mercury (DGM) and radon gas (222Rn) in seawater in the Mediterranean Basin have been measured. Preliminary results indicate similar trends of both gases at some locations, but not at others. Several factors should be considered in order to better understand the relationship between levels of these two gases. The profiles should be interpreted on the basis of the different mechanisms of origin and movement of the gases, as well as their different properties.

Key words: dissolved gaseous mercury (DGM), radon (222Rn), seawater profiles, correlation

Introduction

The main goal of the MERCYMS (Mercury) project is to improve our knowledge of mercury cycling in the Mediterranean Sea. Other constituents of the seawater, which either affect the level of mercury or only accompany it are of interest because they may help better understand the transport and speciation of mercury. Thus, dissolved gaseous mercury (DGM, Hg⁰) and radioactive gas radon (²²²Rn) may both originate in the bottom sediment and be transported to the water surface by carrier gases. Therefore a relationship between their profiles in the water column could be expected.

In oceanic waters, 10-30 % of the Hg may be present in the dissolved gaseous form. While reduction of Hg(II) by aquatic microorganisms was believed to be the predominant source of DGM, recent studies show that photoreduction of Hg may be another important source. [1] Most surface waters are

supersaturated with DGM relative to the atmosphere and therefore elemental Hg is readily lost from water to the atmosphere.^[2]

Radon is produced by radioactive decay of radium (²²⁶Ra) in bottom sediments and of dissolved radium, and may originate from submarine springs and release of fresh waters generally having much higher radon levels than the sea water. [3] In contrast to mercury, radon is a noble gas and does not undergo chemical changes on its travel to the surface, but alpha - decays into its short-lived progeny.

In seawater, Mn – ²²²Rn, CH₄ – ²²²Rn and ³He – ²²²Rn correlations have been obtained when studying heat and chemical flux from the sea bottom. ^[4] Hg – Rn correlation has been studied in soil gas at eruptive sites ^[5] but not in seawater. Results on DGM and ²²²Rn in seawater taken and analysed during the Urania cruises in the Mediterranean Sea in summer 2003 and spring 2004 are re-

ported, and the correlation between the profiles is discussed.

EXPERIMENTAL

Sampling locations are shown in Fig. 1 and their coordinates given in Table 1. Water samples were taken by the rosette at the de-

sired depths. In the laboratory, DGM was purged by $\rm N_2$ from 0.5 dm³ water samples and collected in a gold trap, which was then transferred to a CV AFS analyser system. Radon was degassed from 1.7 dm³ water samples by bubbling in a closed loop and then measured with an 0.7 dm³ evacuated alpha scintillation cell on an alpha scintillation counter.

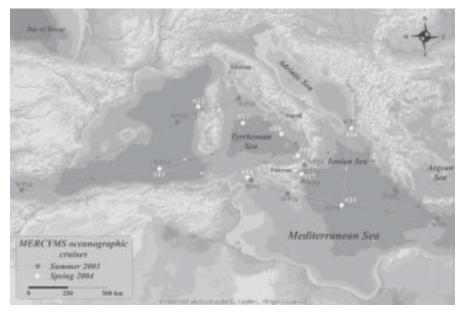


Figure 1. Sampling locations.

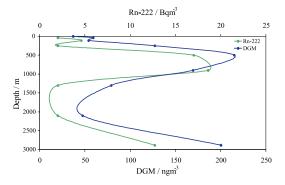
RESULTS AND DISCUSSION

Table 1 summarises the experimental parameters and shows the DGM – ²²²Rn correlation. Four typical depth profiles are shown in Figs. 2-5. The following DGM – ²²²Rn concentration relationship has been observed: 3 strong (Figs. 2 and 3), 3 moderate and 4 weak positive correlations, and 1 strong and 4 weak (Figs. 4 and 5) negative correlations. No final conclusions can be drawn on

the basis of these preliminary results, but we do hope that after our DGM and ²²²Rn data have been further elaborated together with other data collected during the two MERCYMS cruises, we will be able to better interpret the DGM and ²²²Rn profiles observed. In addition, water samples were collected and will be analysed for ²²⁶Ra, in order to provide additional information on ²²²Rn balance. ^[6]

Location	Latitude (N)	Longitude (E)	Date of measurement	Number of sampling	Bottom m	DGM – ²²² Rn correlation
	(14)	(L)	measurement	levels	111	coefficient
WPT 3	37°16'	11°53'	6.8.2003	7	95	- 0.17
WPT 4	35°45'	17°55'	8.8.2003	10	4060	0.35
WPT 5	34°19'	24°20'	10.8.2003	9	2235	- 0.25
WPT 11	37°37'	15°16'	15.8.2003	8	728	0.73
WPT 12	40°34'	14°17'	16.8.2003	5	940	- 0.92
ST. 1	40°29'	11°18'	19.3.2004	8	2883	0.85
ST. 2	41°25'	7°59'	21.3.2004	10	2600	- 0.03
ST. 3	37°52'	5°21'	25.3.2004	9	2816	0.08
ST. 4	37°29'	11°34'	27.3.2004	10	945	0.30
ST. 5	35°45'	17°55'	29.3.2004	12	4040	0.49
ST. 6	39°59'	19°00'	30.3.2004	10	919	0.09
ST. 7	37°37'	15°15'	1.4.2004	11	683	0.41
ST. 8a	38°39'	15°05'	2.4.2004	6	40	- 0.25
ST. 8b	38°39'	15°06'	2.4.2004	7	73	0.86
ST. 9	39°55'	14°00'	3.4.2004	10	2380	0.40

Table 1. Description of sampling locations and DGM – ²²²Rn correlation coefficients.



Rn-222 / Bqm³

0 2 4 6 8 10

100

200

E 300

600

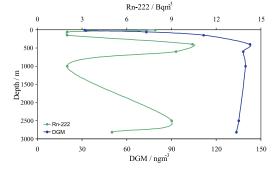
700

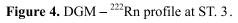
800

DGM / ngm³

Figure 2. DGM - ²²²Rn profile at ST. 1.

Figure 3. DGM - ²²²Rn profile at WPT. 11.





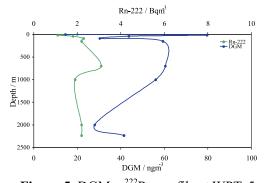


Figure 5. DGM - ²²²Rn profile at WPT. 5.

CONCLUSIONS

Our preliminary results indicate that at some locations, DGM and ²²²Rn concentrations in the seawater column are well correlated, while such a correlation has not been found at others. No general rules have been eluci-

dated either from data obtained for various locations in the same season, or from comparing data obtained in summer and spring at the same station. Final conclusion will be possible after the data have been further elaborated and results of ²²⁶Ra analyses will be included in the interpretation.

REFERENCES

- [1] Costa, M. & Liss, P. (2000): Photoreduction and evaluation of mercury from seawater; *Science of the Total Environment* 261, pp. 125-135.
- [2] HORVAT, M., KOTNIK, J., LOGAR, M., FAJON, V., ZVONARIĆ, T. & PIRRONE, N. (2003): Speciation of mercury in surface and deep-sea waters in the Mediterranean Sea; *Atmospheric Environment* 37, pp. S93-S108.
- [3] CORBETT, D. R., DILLON, K., BURNETT, W. & CHANTON, J. (2000): Estimating the groundwater contribution into Florida Bay via natural tracers ²²²Rn and CH₄: Limnology and Oceanography 45, pp. 1546-1557.
- [4] ROSENBERG, D. N., LUPTON, J. E., KADKO, D., COLLIER, R., LILLEY, M. D. & PAK, H. (1988): Estimation of heat and chemical fluxes from a seafloor hydrothermal vent field using radon measurements; *Nature* 334, pp. 604-607.
- [5] Thomas, D. M. (1986): Geothermal resources in Hawaii; *Geothermics* 15, pp. 435-514.
- [6] SARMIENTO, J. L., BROECKER, W. S. & BIACAYE, P. E. (1978): Excess bottom ²²²Rn distribution in deep ocean passages; *Journal of Geophysical Re*search 83, pp. 5069-5076.